



Assimilation of GPS radio occultation measurements at the Met Office

by Michael Rennie

Presentation to **GRAS SAF Workshop on Applications of GPS radio occultation Measurements**, 16 -18 June 2008

© Crown copyright Met Office



Talk outline

1. Met Office assimilation system / observation operator
2. Status of RO in operations
3. Recent impact studies
4. Future developments



Met Office system: brief summary

- NWP model is the Unified forecast model (UM):
 - Non-hydrostatic equations
 - Height as the vertical co-ordinate.
 - Charney-Philips grid-staggering in the vertical.
 - Terrain-following near the surface.
- Variational data assimilation system (VAR):
 - Incremental 4D-Var.
 - Uses perturbation forecast (PF) model to map background error info to the time of the observations
 - PF has simplified linearised physics, rather than direct tangent linear/adjoint of non-linear UM.

RO assimilation method within VAR:

- Given state vector \mathbf{x}_1 on the 1st iteration:

- Forward model N using **non-linear operator**: $\mathbf{y}_1 = H(\mathbf{x}_1)$

- Calculate local gradient i.e. jacobian, sometimes called \mathbf{K} matrix: $\frac{\partial \mathbf{y}_1}{\partial \mathbf{x}_1} = \frac{\partial H(\mathbf{x}_1)}{\partial \mathbf{x}_1} = \mathbf{K}$

- Store $\mathbf{x}_1, \mathbf{y}_1$ and the \mathbf{K} matrix.
- On subsequent iterations given incremented state vector \mathbf{x}_n :
 - apply the **tangent-linear** approximation to estimate $\mathbf{y}(\mathbf{x}_n)$:

$$\mathbf{y}_n = \mathbf{y}_1 + \mathbf{K}(\mathbf{x}_n - \mathbf{x}_1)$$

- Use \mathbf{y}_n and \mathbf{K}^T matrix to calculate on n^{th} iteration:
 - Observation cost function (J_{obs}):

$$(J_{obs})_n = \frac{1}{2}(\mathbf{y}_n - \mathbf{y}_{obs})^T \mathbf{R}^{-1}(\mathbf{y}_n - \mathbf{y}_{obs})$$

- Gradient of J_{obs} wrt \mathbf{x} :

$$\frac{\partial (J_{obs})_n}{\partial \mathbf{x}_n} = \left(\frac{\partial \mathbf{y}_1}{\partial \mathbf{x}_1} \right)^T \mathbf{R}^{-1}(\mathbf{y}_n - \mathbf{y}_{obs}) = \mathbf{K}^T \mathbf{R}^{-1}(\mathbf{y}_n - \mathbf{y}_{obs})$$

- Total GPSRO J_{obs} and gradient information used with contributions from other observation data in the minimisation problem to produce an updated \mathbf{x} .

1D Refractivity operator summary

1. Interpolate model column data to occ time and location: to give x

GPSRO x is pressure (P) on ρ levels & relative humidity (RH) on θ levels

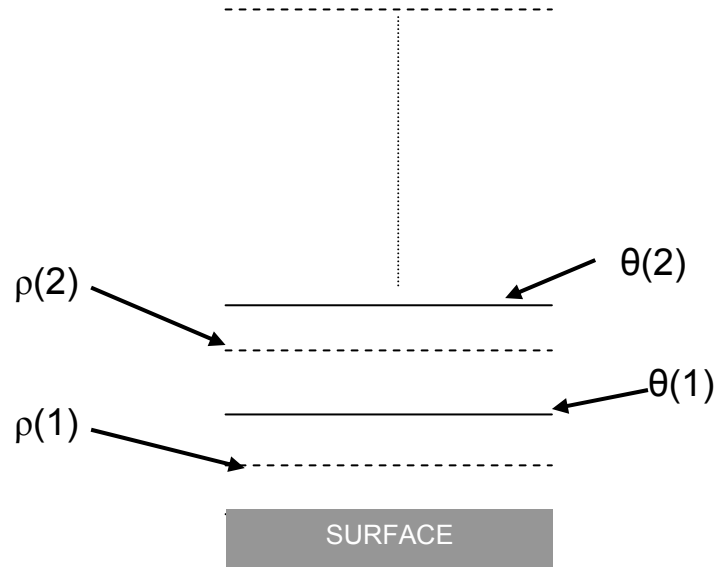
2. Interpolate Exner from ρ to θ levels to get P on θ levels

3. Calculate layer mean *virtual* temperature on θ levels

4. Calculate layer mean temperature on θ levels, using RH

5. N calculated on θ levels using Smith-Weintraub formula

6. N interpolated to obs heights





1D refractivity operator strengths/weaknesses

- Strengths:
 - Simple and quick.
 - No extrapolation above model top, as required for *BA*.
- Weaknesses:
 - A priori data introduced high up ($>\sim 25$ km) from climatology in *N* data.
 - **R** matrix more complicated for *N* than *BA*?
- Future updates:
 - Use *q* instead of RH.
 - Adjust code for *BA* assimilation.
 - Met Office system not yet capable of incorporating 2D operators.

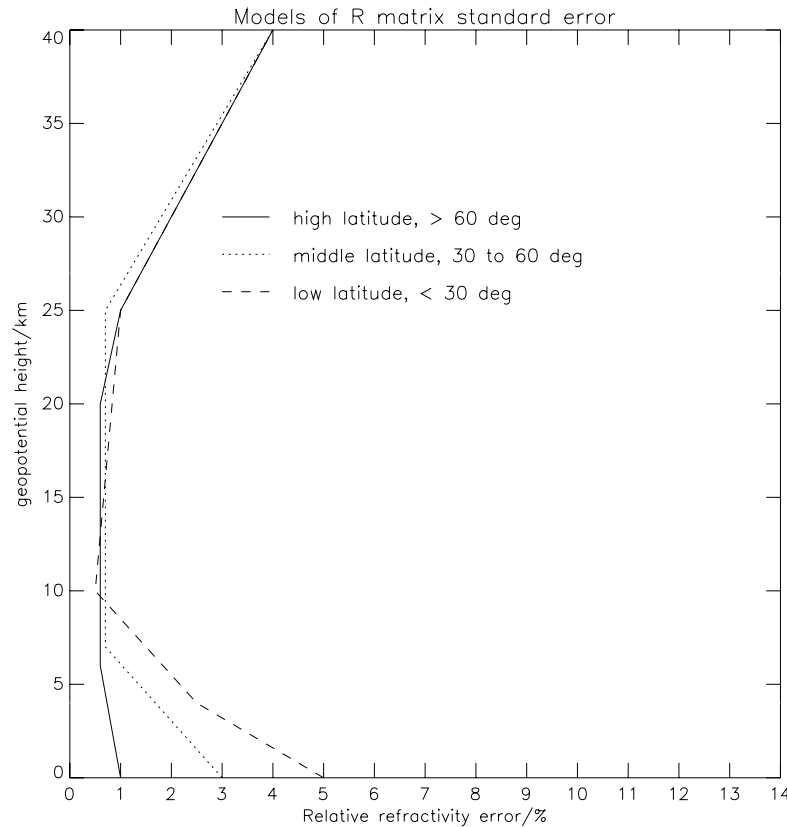


GPSRO used in global model

- UM:
 - Ran at N320L50 i.e. ~40 km mid lat horizontal resolution, 50 vertical levels, top ~ 0.1 hPa (~63 km).
 - Forecasts out to 6 days.
- VAR:
 - 6 hour assimilation window.
 - First iteration non-linear using N320 3, 6, 9 hour forecast background information.
 - Subsequent tangent-linear iterations use increments to model columns.
 - Non-linear iteration is ran on every 10th iteration.
 - Other significant data types assimilated: Sonde, IASI, AIRS, ATOVS, Aircraft, Satwind, Scatwind, Surface, SSMI, SSMIS

Global model: GPSRO specific

- Use **R** matrix for low, middle and high latitudes.
 - Based on $(O-B)/B$ std dev using COSMIC.

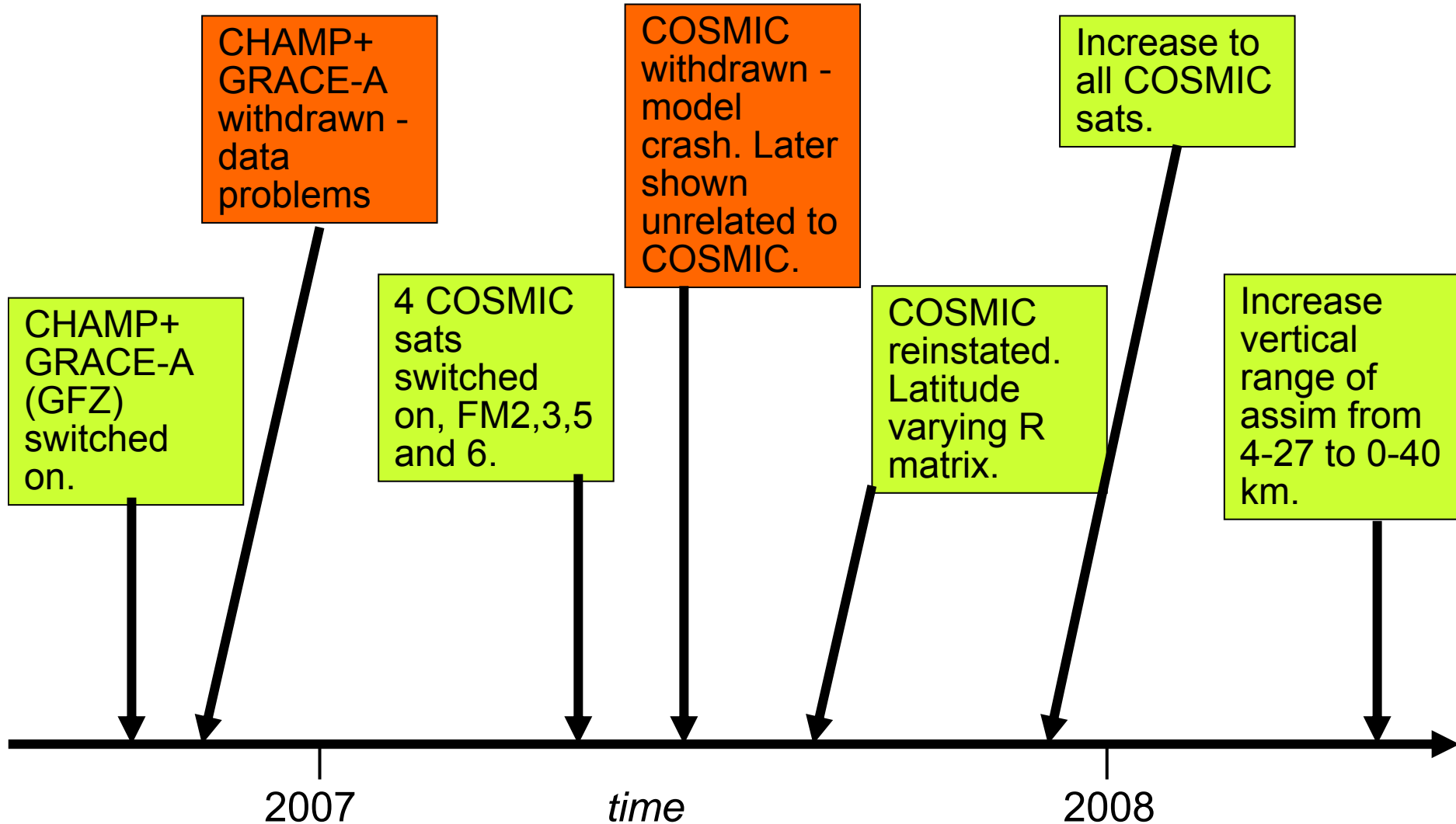


$$R_{ij} = \sigma_i \sigma_j \exp\left(\frac{-|z_i - z_j|}{H}\right)$$

- Assume an exponentially decaying vertical correlation model with a scale length of 3.3 km.
- QC of N data based on output of a 1D-Var.



Global model operational status

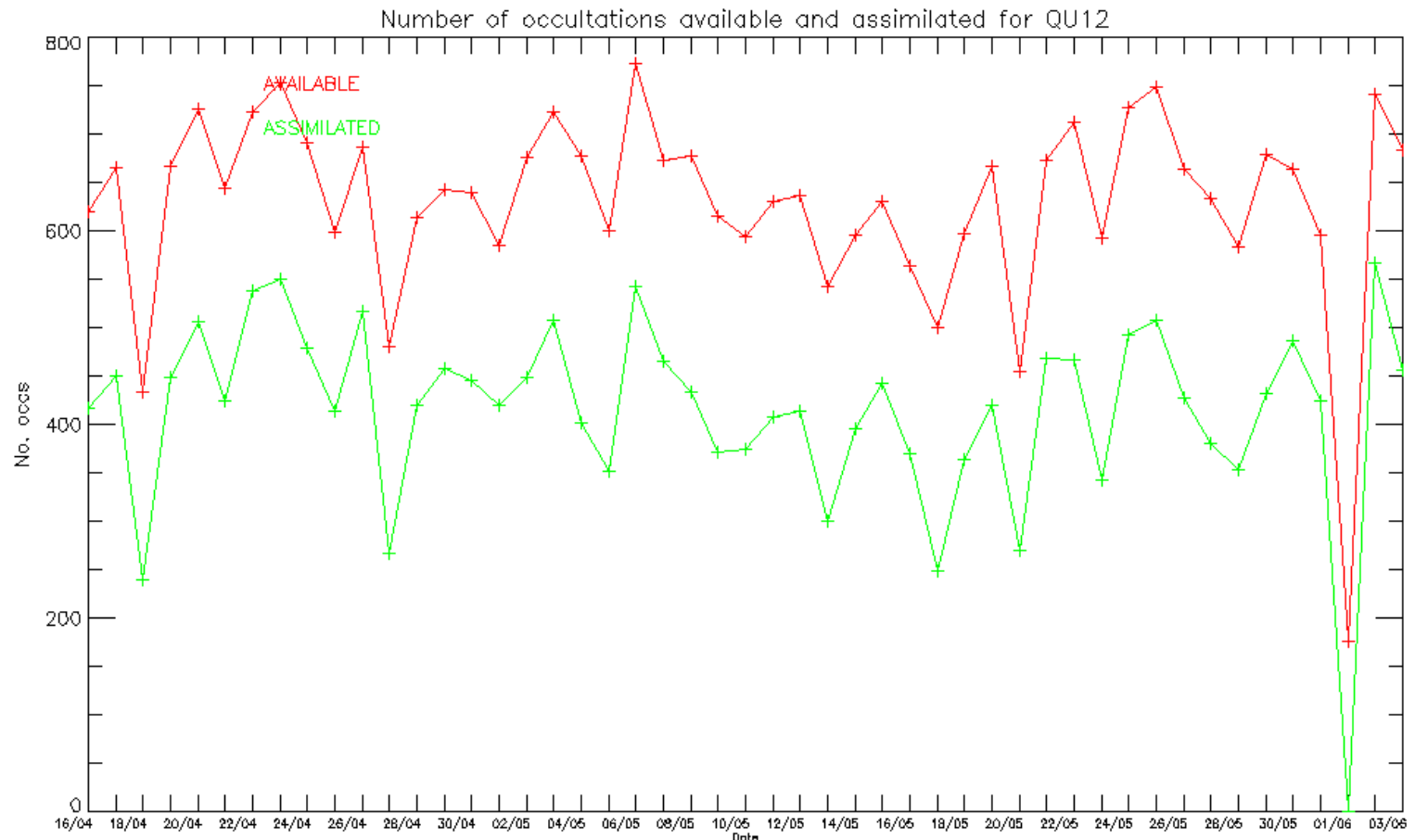




Number of N profiles assimilated in global model

Update run to produce better background for next run

- **Assimilated** = COSMIC data
- **Available** = COSMIC + GRAS+GRACE -A+CHAMP

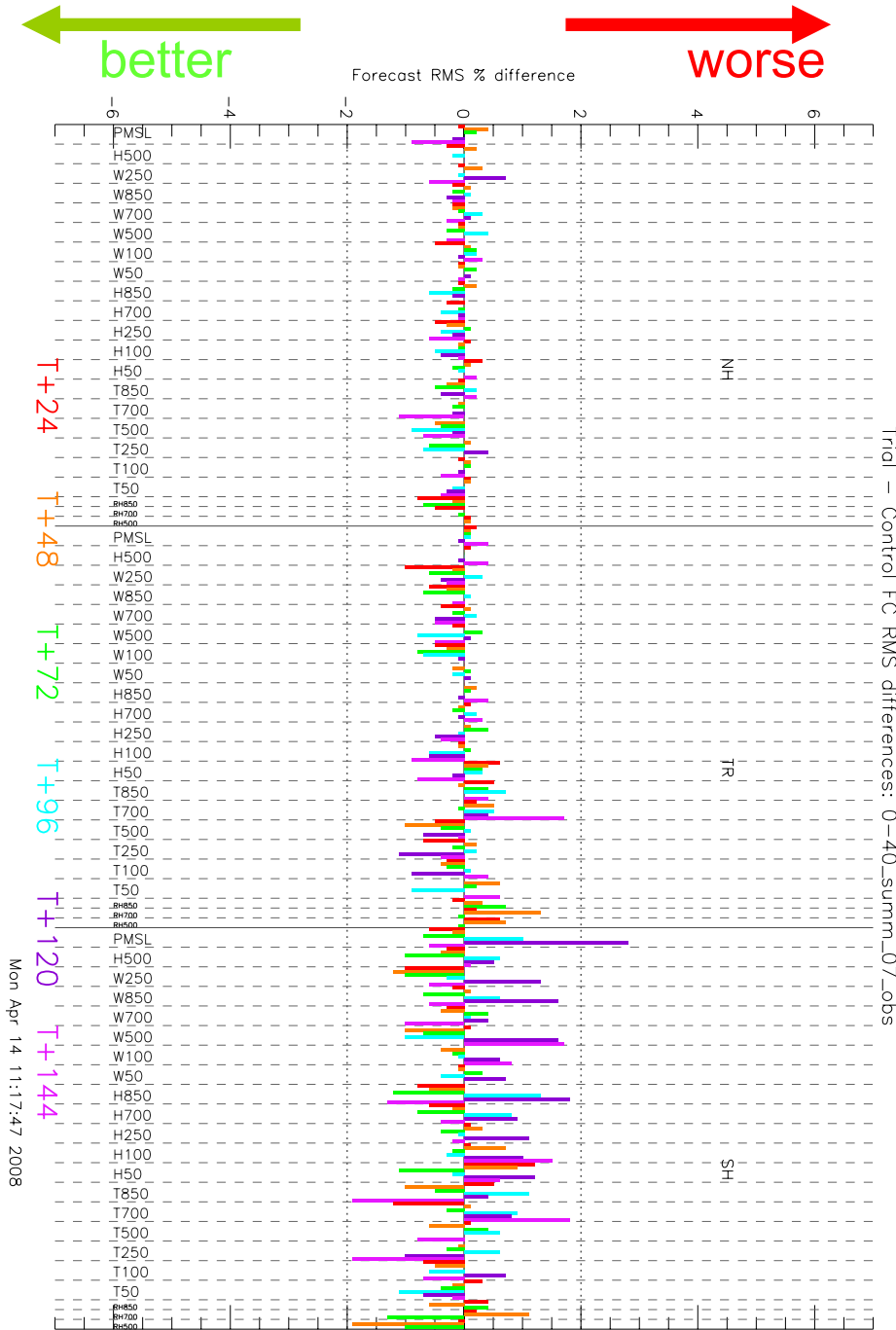


Wed Jun 4 04:46:37 2008



Impact studies: increasing vertical range

- Increasing the vertical range of assimilation from 4-27 km to 0-40 km gave a small benefit to:
 - Tropospheric relative humidity in extratropics.
 - Winds - highly valued by customers.
 - Stratosphere model bias.
- Routine verification against sondes and analyses is only up to 50 hPa (~21 km).



Forecast RMS % diff. against **obs** (mainly radiosondes)

Control= 4-27 km

Exp=0-40 km
Jun 2007



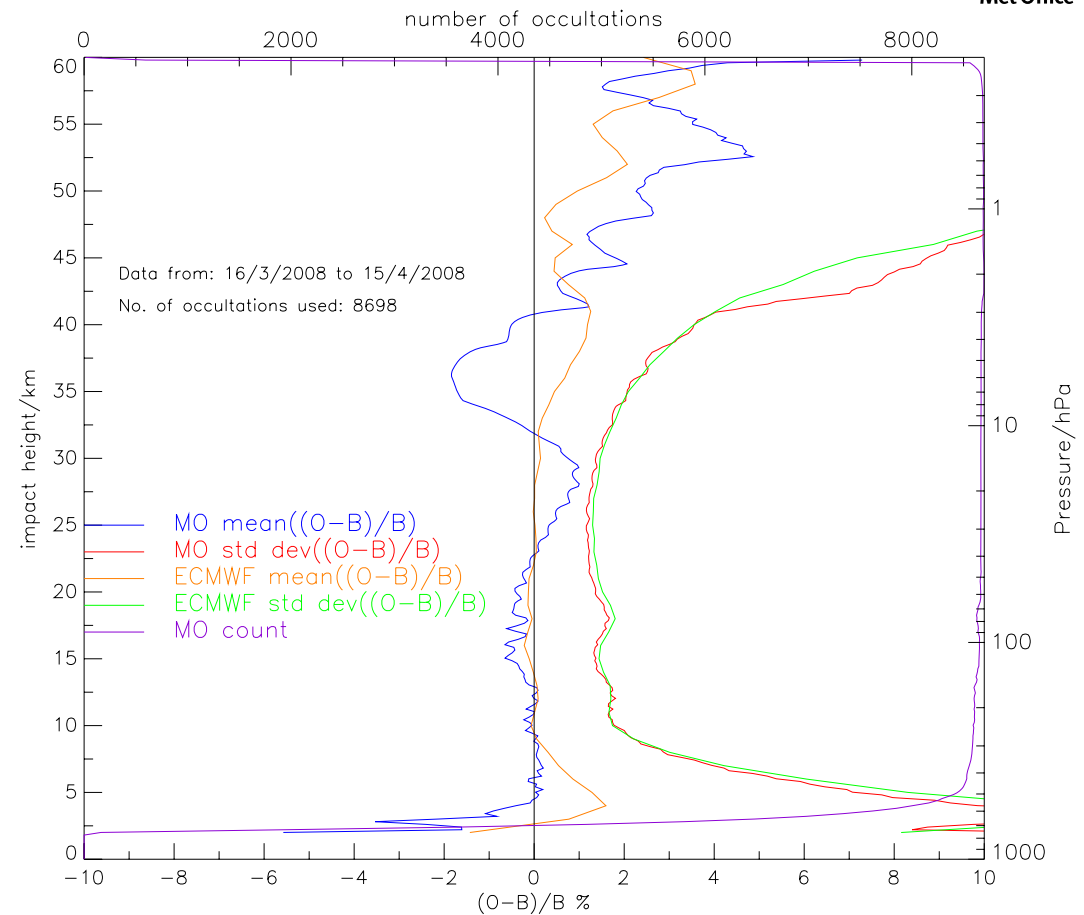
Stratosphere global model bias compared to COSMIC

UCAR processed Cosmic1 data, global

Bending angle bias and standard deviation plots



- $BA (O-B)/B$ stats show a distinct 'S' shape bias at > 50 hPa (~ 22 km, around model level 35).
- ECMWF stats shown for comparison: has its own biases



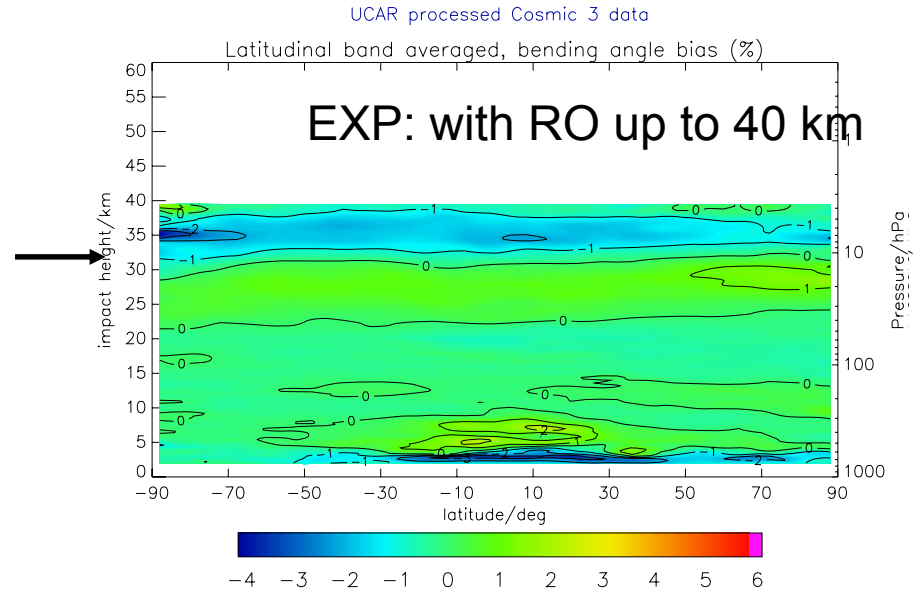
See latest GRAS SAF monitoring: <http://monitoring.grassaf.org>



Impact on bias of RO up to 40 km

- Stratospheric bias reduced by *N* assimilation up to 40 km.
- Plot uses BA before statistical optimisation.

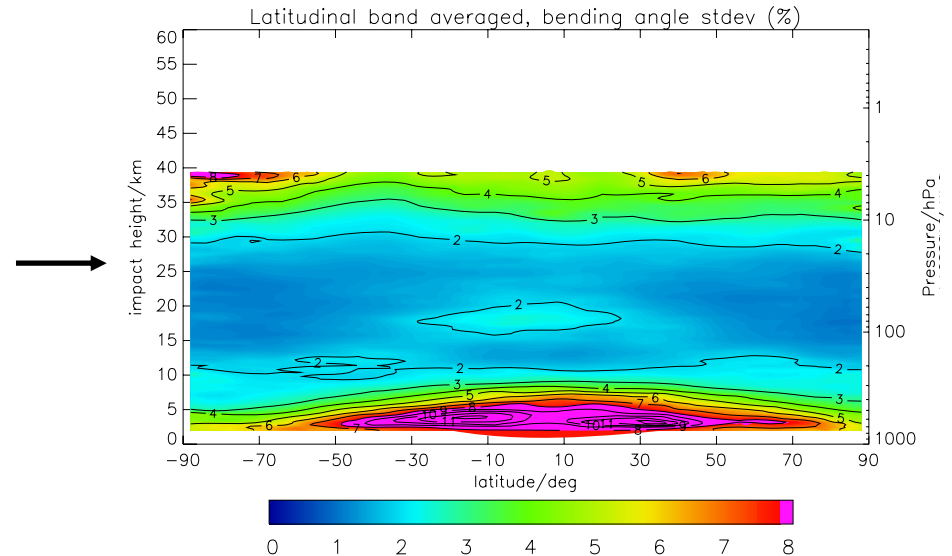
Mean
(O-B)/B



Data from: 25/5/2007 to 1/6/2007

Plotted at: 14:56 28-Nov-2007

St. Dev.
(O-B)/B

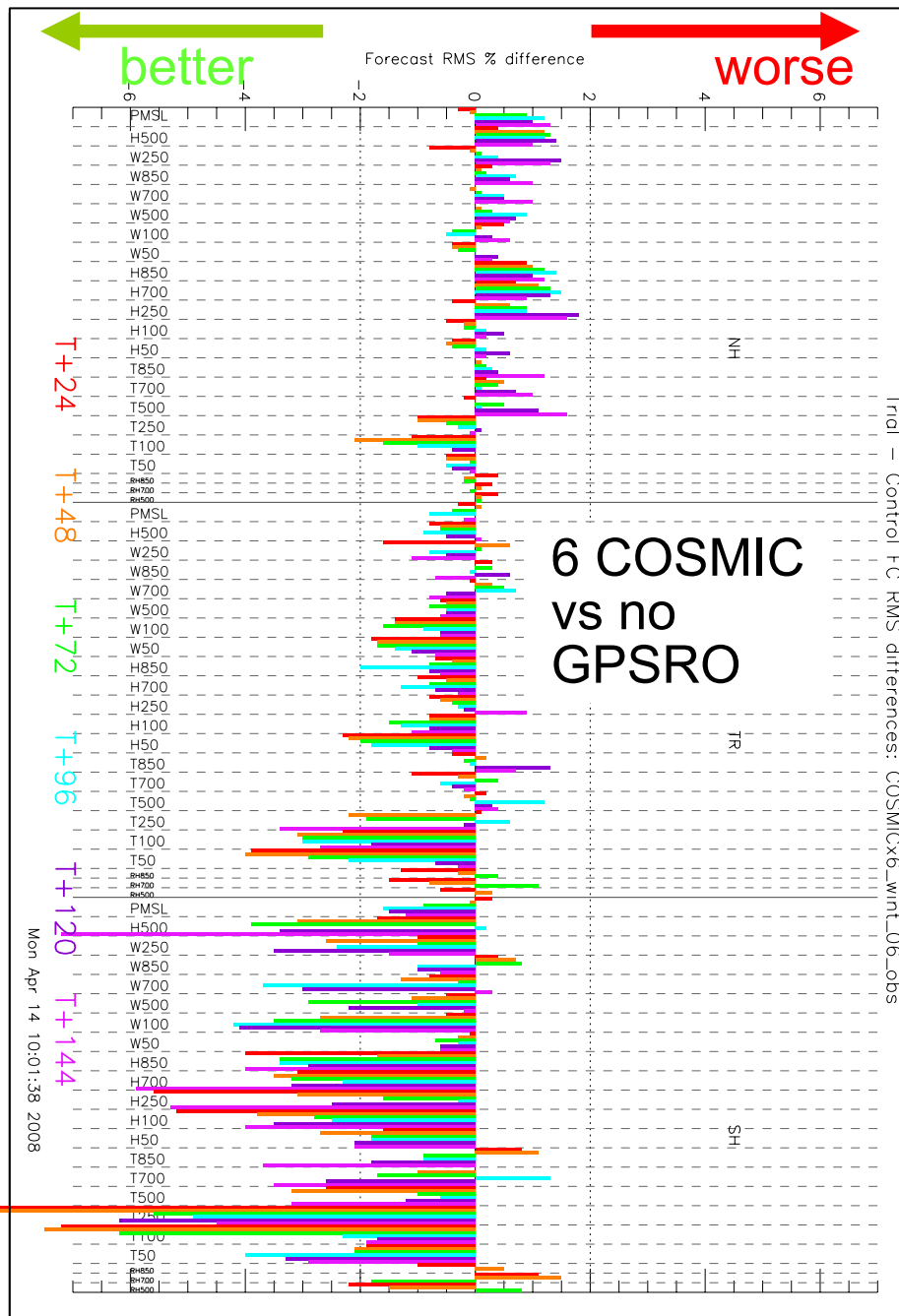




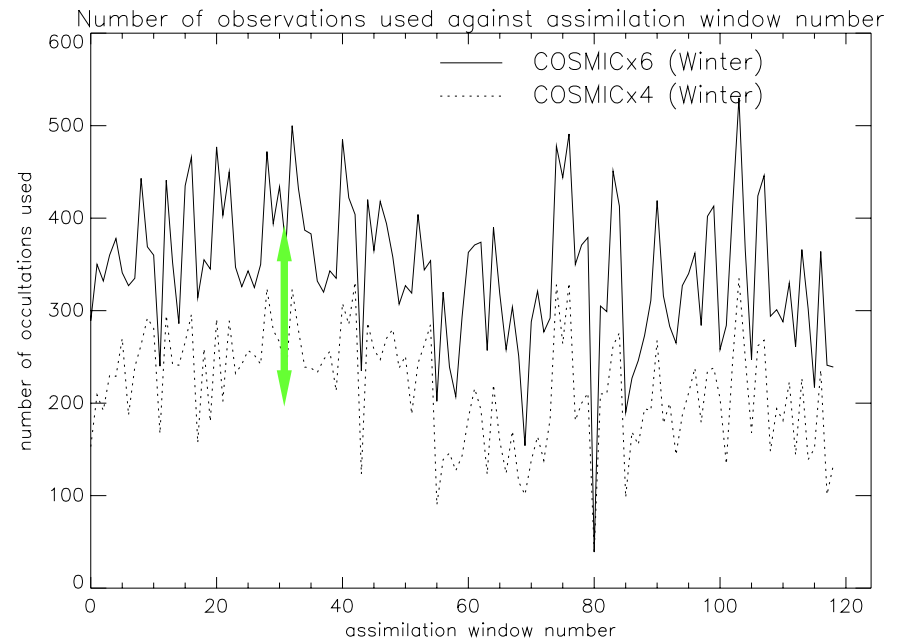
Impact of using more RO data

- More data seen to increase magnitude of the impact:
 - Dec 2006, going from 4 to all 6 COSMIC sats.
 - Jun 2006, going from 4 COSMIC to (6 COSMIC +CHAMP+GRACE-A).
 - Jan 2008, GRACE-A and CHAMP (GFZ) on top of COSMIC. Small improvements in geopotential height.
- Would be interesting to run experiment using incrementally more data - how saturated with RO are we?

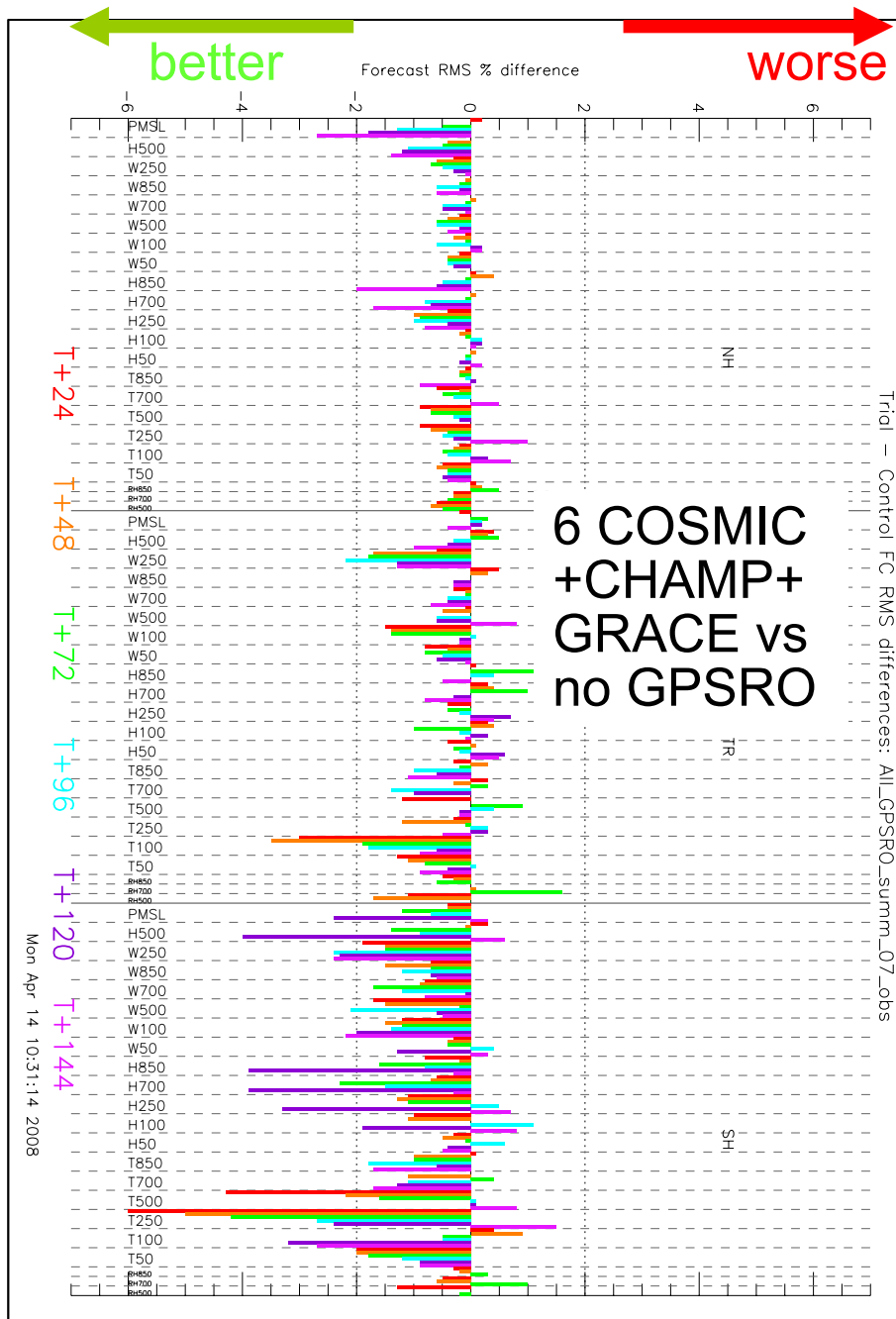
Forecast RMS % diff. against obs, Dec 2006 trials



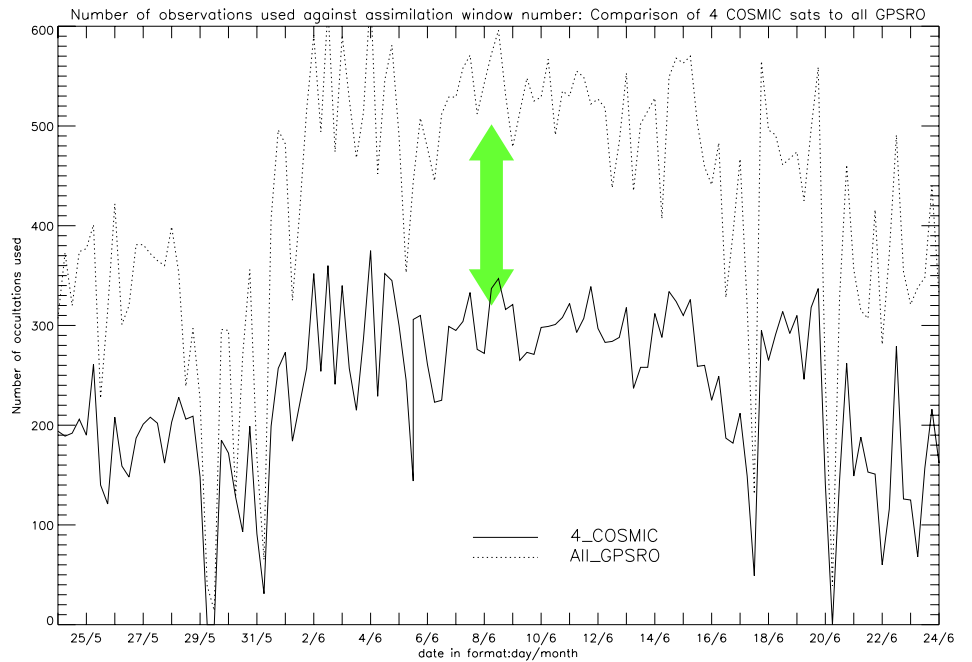
Increase in no. of occultations:



Forecast RMS % diff. against obs, Jun 2007 trials



Increase in no. of occultations:



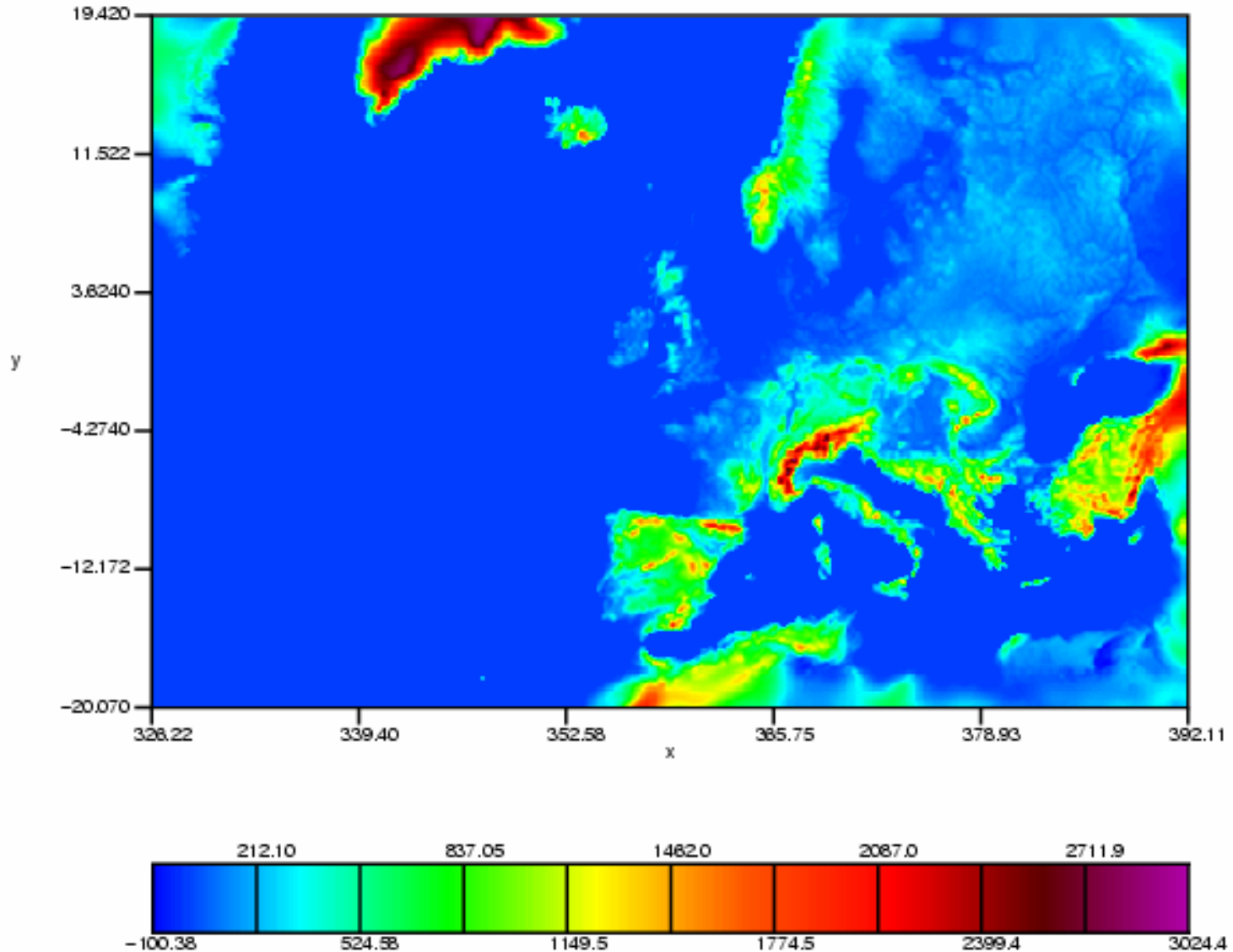


GPSRO in limited-area models

- Ob types implemented into the global model then go into limited-area models with relatively little testing.
- Concerns higher horizontal resolution - problems using 1D N operator.
- Ran test of N assimilation using 1D operator.



NAE (North Atlantic and European) area





NAE test setup

- 24 km resolution (half operational res. to reduce time).
- 38 vertical levels, ~40 km model top. 4D-Var.
- Typically around 20-30 occs assimilated per cycle. All COSMIC +CHAMP+GRACE-A, using 0-40 km vertical range.
- 20 day period of testing from 24/04/08 to 26/05/08 (with some gaps)



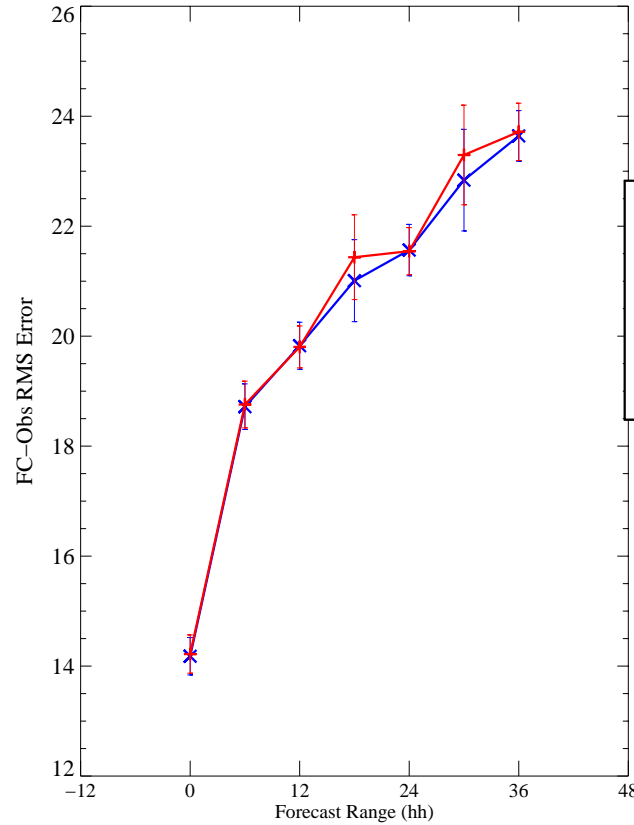
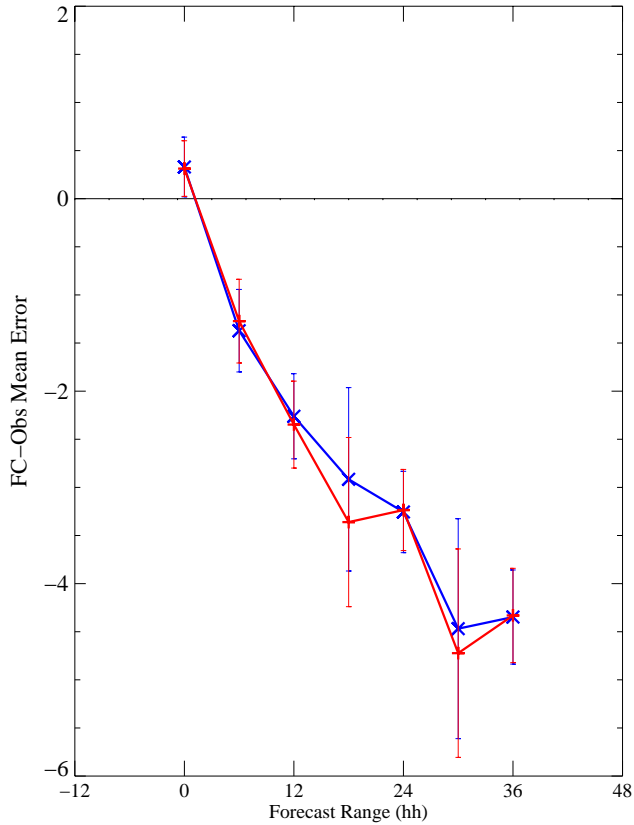
NAE test results

- Verified using limited area **NWP index**. A Met Office score system based on comparisons to observed fields useful in limited-area model forecasting:
 - Surface visibility, 6 hr precipitation accumulation, total cloud amount, cloud based height (3/8 Cover), surface temp. and surface wind.
- Saw a small **overall improvement**. Particularly in surface visibility.
- NAE area NWP index increased by +0.13 %, i.e. slightly positive. UK area NWP index +1.23%, although significance in question over limited area and short period.

Some GPSRO in NAE verification



Cases: + Control x GPSRO in NAE



700 hPa
RH (%) vs
sondes.
NAE area.

CONTROL

GPSRO

Mean error

RMS error

68% error bars calculated using $S/(n-1)^{1/2}$



Future updates/plans

- **operational changes**, planned for July 2008.
 - **Global model:** use of MetOp GRAS (10-30 km vertical range), CHAMP+GRACE-A (GFZ) on top of COSMIC.
 - **NAE model:** use of COSMIC, CHAMP and GRACE-A (0-40 km):
- **Further tuning** of system:
 - Obs errors and correlations.
 - Vertical ranges.
- **Experiment** with BA assimilation



Any questions?